

Estimating Specific Features of the Optical Property Variability in the Black Sea Waters Using the Data of *SeaWiFS* and *MODIS* Satellite Instruments

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Problem of quantitative interpretation of the Black Sea observations provided by the *SeaWiFS* and *MODIS* satellite instruments is considered in the terms of spectral dependencies of the seawater optically active components. For this purpose, the results of standard data (accumulated in the NASA archive) processed by the *GIOP* (Generalized ocean color inversion model for retrieving marine Inherent Optical Properties) complex method are analyzed. As these results often contain significant distortions related to the atmospheric disturbances, selection of reliable test data implies the following requirement: large cloudless areas of the sea should contain no sudden chaotic spatial-temporal fluctuations of all the defined products and no false local correlations between the fields of the atmospheric and seawater parameters. Besides, imposed are the conditions for sufficiently accurate model reproduction of the empirical spectra of the sea surface reflectance and coincidence of the results obtained from the *SeaWiFS* and *MODIS* instruments. Application of the *GIOP* method permits to analyze the features of variations and the relative role of the light absorption components in the upper water layer associated with phytoplankton and the dissolved yellow substance. In the deep-water part of the Black Sea in summer, yellow substance makes the main contribution to absorption, and during the summer-autumn transition period, approximately equal growth of both the phytoplankton and the yellow substance absorptions is observed. Having been compared, the features of the Black Sea and the Equatorial Pacific waters are represented as an example. In contrast to the Black Sea, phytoplankton in the Equatorial Pacific is a dominating factor, whereas the yellow substance content remains almost unchanged.

Keywords: the Black Sea, the Pacific Ocean, satellite observations, spectral dependencies, optical characteristics, *MODIS*, *SeaWiFS*, *GIOP*, phytoplankton, yellow substance.

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Introduction

The information obtained by *SeaWiFS*, *MODIS* optical instruments and others operating in space can be used to control and study the various processes of the marine water area state variability. The results of regular global satellite observations and products of their processing are accumulated on an on-going basis in special archives with free access via the Internet. The equipment installed on the

satellite registers accurately the outgoing radiation of various wavelengths of the visible spectrum range, which in a complex manner depends on a multitude of variable parameters of the atmosphere and the upper sea layer. These include the coefficients of light absorption $a(\lambda)$ and backscattering $b_b(\lambda)$ in water (λ is the wavelength of light), chlorophyll a concentration, optical thickness of the aerosol component of the atmosphere, etc.

One of the directions of using satellite measurements is to determine the numerical values of these parameters. For this purpose, different models, methods and algorithms can be applied. Simplified approaches convenient for carrying out mass calculations are most often used. In these, in an explicit or implicit form, it is assumed that the number of independent factors the results of satellite measurements depend on is substantially limited. Such methods include, for example, those described in [1–5]. Such methods cannot provide high efficiency in interpreting observational data in a wide variety of conditions that vary from region to region, in different seasons and so on.

The complex methods designed for more accurate accounting of all the main effects manifested in variations in the optical properties of seawater [6–8] are also known. In recent years, the NASA operating system for satellite data processing has been supplemented with a procedure implementing one of such complex methods – *Generalized ocean color inversion model for retrieving marine Inherent Optical Properties (GIOP)* [9]. Accordingly, a wider set of parameters including $a(\lambda)$ and $b_b(\lambda)$ also includes the light absorption in the sea due to phytoplankton $a_{ph}(\lambda)$ and yellow substance $a_{dg}(\lambda)$ (in combination with detritus) content. NASA archives have added the *GIOP* application results for the accumulated earlier global observation arrays for the entire operation time period of *SeaWiFS* and *MODIS* instruments in space.

GIOP is based on physical models that describe in a parametric form the dependence of the optical properties of the sea on the light wavelength and the content of the main optically active impurities in water. These models are obviously more or less approximate in nature. Therefore, it is required to verify the *GIOP* operating capacity, taking into account the specific features of the optical properties of the Black Sea water. For this purpose, it is useful to consider in detail some of the most characteristic examples that reflect the specific features of the Black Sea waters. Similar studies were earlier performed in [7, 10] using other methods. Some results of the *GIOP* application are described in [11], but there a somewhat different problem was considered.

The Black Sea

The initial data in the calculations of the seawater characteristics by the *GIOP* method are the spectral values of the sea surface reflectance $R_{rs}(\lambda)$ for all measurement channels of the visible range, which are obtained at one of the first stages of the satellite observations processing in atmospheric correction. As is known, for the Black Sea the results of $R_{rs}(\lambda)$ calculation can often contain significant errors. The main origin for them is the difference between the optical properties of a real atmospheric aerosol from those models that are incorporated

into the NASA-developed data processing system for global satellite observations. In addition, the accuracy of $R_{rs}(\lambda)$ determination is influenced by factors such as the sun glint, the errors in setting the calibration parameters of the instrument, the increased sensitivity of the sensors to changes in the recorded radiation polarization, etc.

Consequently, to obtain reliable conclusions it is necessary to avoid the use of distorted data and limit to analyzing a relatively small number of examples with the most reliable initial values of $R_{rs}(\lambda)$. For their selection, the criteria used earlier in [10–13] can be applied. According to these criteria, it is required that within the large cloudless sea areas there are no sharp chaotic spatio-temporal fluctuations of all determined parameters (including atmospheric ones) and false local correlations between the products of calculations of the atmosphere and sea water characteristics. In addition, a sufficiently accurate model reproduction of $R_{rs}(\lambda)$ empirical spectra and the coincidence of the results obtained from SeaWiFS and MODIS instruments are needed. Below the results of the analysis of test data obtained using these criteria are considered.

Fig. 1 shows charts of the spatial distribution of chlorophyll *a* concentration (denoted by C_a) on 11.08.2006 and 04.10.2006. Here, the results of C_a calculation by the traditional simplified method [1] are used. As is known, in the Black Sea this method can lead to errors, nevertheless, the charts in Fig. 1 indirectly represent the optical properties of water in these days.

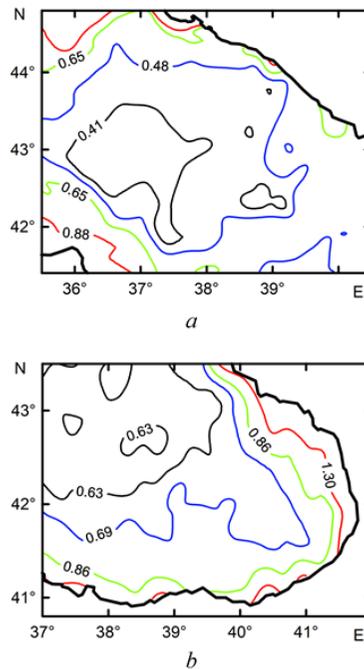


Fig. 1. Charts of spatial distribution of value C_a , $\text{mg}\cdot\text{m}^{-3}$ in the eastern Black Sea on 11.08.2006 (a) and 04.10.2006 (b)

In the deep-water part of the sea on 4.10.2006 within a large area the C_a field was almost homogeneous and in the east, as the coast approached, the C_a growth took place. It is essential that on this day almost the entire eastern part of the Black Sea was free from cloud cover and provided with results of practically simultaneous surveys by *MODIS* and *SeaWiFS* instruments. For the test area considered below, the time difference between the two instruments does not exceed 20 minutes.

Fig. 2 shows the values of the $a(443)$ and $a_{ph}(443)$ calculated by the *GIOP* method, selected from the band located about $42,0^\circ\text{N}$ and having a width of $0,25^\circ$ in latitude. In these graphs, the *MODIS* data in the *Level-3m* format obtained on 4.10.2006 were used.

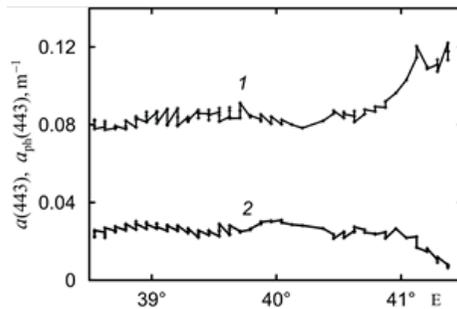


Fig. 2. Variations of the absorption parameters obtained on 04.10.2006 within the narrow strip close to $42,0^\circ\text{N}$: $a(443)$ is curve 1 and $a_{ph}(443)$ is curve 2

In Fig. 2 on the area to the west from $\sim 41^\circ\text{E}$ $a(443)$ and $a_{ph}(443)$ values do not almost change. But approaching the coast in the east, $a(443)$ increases in a natural way, but the results of the $a_{ph}(443)$ determination show a significant decrease. The chart in Fig. 1 shows an increase in the chlorophyll a concentration in the same region, and therefore an increase in $a_{ph}(443)$ should also be expected. It is noteworthy that, there is no such decrease in $a_{ph}(443)$ in the *SeaWiFS* data. These contradictions are a clear example of the manifestation of distortions in the *GIOP* method application results.

Such distortions are often found in the analysis of satellite data processing products using similar to *GIOP* complex methods, aimed at simultaneous determination of many unknown seawater parameters based on measurements in a complete set of spectral channels of instruments operating in space. This is the general property of such (синонимичные конструкции) problems, consisting in the ambiguity and instability of their solutions with respect to various inaccuracies in the model and in the initial data. A detailed study of this problem is beyond the scope of the present article, some of its aspects are considered in [2, 8, 10, 14].

The *Level-3m* format data used here and below are the final processing products interpolated to the regular grid nodes, for which all necessary transformation operations are carried out, including automatic rejection of hardly usable samples for a number of formal criteria. The analysis in Fig. 2 and many other similar examples of the Black Sea observations suggests that in the *GIOP* method implementation such a rejection does not completely exclude the distorted

data. The aforementioned criteria are more effective. At the same time, despite the fact that *MODIS* and *SeaWiFS* instruments often display various serious distortions in the Black Sea observations, it remains possible to obtain useful information by careful selection of observational data in the most favorable situations.

In the example of the eastern Black Sea survey on 4.10.2006, a quite favorable area is located near 39°E within the data sampling used in Fig. 2. Fig. 3 and Tab. 1 shows the results of applying the *GIOP* to *MODIS* data for one of the grid nodes in this area. Here, the results of the Black Sea survey data processing obtained in the same way on 11.08.2006 are also given.

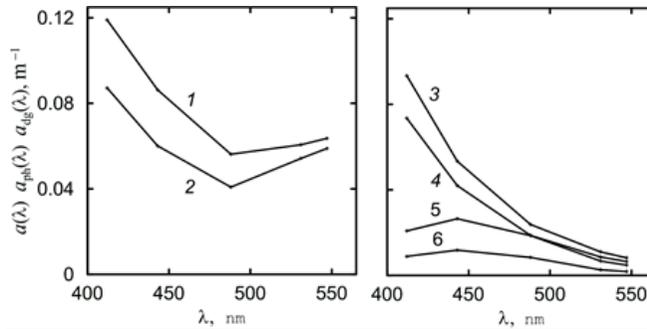


Fig. 3. Results of determining the absorption spectral characteristics by the *GIOP* method based on the Black Sea surveys performed by the instrument *MODIS* on 04.10.2006 (1, 3, 5) and on 11.08.2006 (2, 4, 6): 1 and 2 are $a(\lambda)$; 3 and 4 are $a_{dg}(\lambda)$; 5 and 6 are $a_{ph}(\lambda)$

Table 1

Results of determining the chlorophyll *a* concentration and the optical parameters of the Black Sea waters using the *GIOP* method

Date	Coordinates		$a_{ph}(443), m^{-1}$	$a_{dg}(443), m^{-1}$	$C_a, mg \cdot m^{-3}$	$C_{ag}, mg \cdot m^{-3}$
	°N	°E				
11.08.2006	42.37	37.71	0.012	0.042	0.42	0.21
4.10.2006	42.21	39.04	0.025	0.055	0.72	0.48

The examples in Fig. 3 and in Tab. 1 satisfy the above criteria, including the condition that the results of calculations of all parameters and products of the *GIOP* application for the two devices coincide. It is also essential that in this case the models used in *GIOP* provide accurate reproduction of the empirical spectral dependences of the sea surface reflectance $R_{rs}(\lambda)$.

These examples are interesting in that they were obtained in the same year in the same sea area and reflect the usual changes in optical properties from summer to autumn. Despite the fact that the data in Fig. 3 and in Tab.1 refer to readings in single nodes of the grid of the *Level-3m* format, these data characterize the properties of extended areas in the deep-water part of the Black Sea.

The *GIOP* method application results allow analyzing the features of the variations and the relative role of the components that affect the light absorption in the upper layer of water. In the transition from summer to autumn, approximately the same increase is observed in each of $a_{ph}(\lambda)$ and $a_{dg}(\lambda)$ values, but the relative changes

of $a_{dg}(\lambda)$ are less pronounced, since this component makes the main contribution to the total absorption. The latter circumstance serves as one of the key sources of errors in the determination of C_a in the Black Sea by satellite measurements.

It should be noted that conclusions similar in meaning were formulated earlier in [7, 15–17] based on the field data analysis and in [2, 10] – in the interpretation of satellite observations. The new results obtained in the present study contain a more detailed analysis of key effects and are of particular interest as they characterize concrete examples of the *GIOP* method application.

According to [18], in the waters of the open part of the Black Sea, the light absorption by detritus is much less than $a_{ph}(\lambda)$. So, it can be assumed that the main role is played by the yellow substance, and in the considered situations in the notation of $a_{dg}(\lambda)$, detritus appears only to follow exactly the notation system adopted in the *GIOP* method description [9].

In addition to estimating the chlorophyll *a* concentration, calculated in the traditional simplified way [1], Tab. 1 also demonstrates the results of its calculations by the *GIOP* method, denoted as C_{ag} . Both these values are determined from the data of satellite surveys by the *MODIS* instrument. It is easy to see that C_a exceeds C_{ag} approximately one and a half to two times. This difference is a direct consequence of the specifics of the optical properties of water in the Black Sea. It was previously established that a simplified method [1] often overestimates C_a in the Black Sea [2, 19, 20], so the results of calculations by the *GIOP* method can be considered to be closer to reality.

This is also evidenced by the fact that the application of a specially developed regional algorithm for the Black Sea in [2] based on measurements on 11.08.2006 and 4.10.2006 leads to estimates of the chlorophyll *a* concentration close to C_{ag} : 0.26 and 0.41 $\text{mg}\cdot\text{m}^{-3}$. In addition, the value of C_{ag} obtained for 11.08.2006 is in full agreement with the usual conditions at this time of year in the deep-water part of the Black Sea. For 4.10.2006, the C_{ag} , determined from satellite measurements, and the corresponding $R_{rs}(\lambda)$, are close to the results of in situ measurements described in [7] carried out on 14.09.2000 at one of the stations during the expedition in the northeastern part of the sea.

The Pacific Ocean

For clearer understanding the *GIOP* method possibilities and evaluation the features of the optical properties of the Black Sea water, it is useful to compare them with the properties of waters of other areas. To this end, Tab. 2 and Fig. 4 gives examples of *SeaWiFS* data processing products specially selected from surveys of the equatorial region of the Pacific Ocean carried out on 29.05.2003 and 13.06.2010. Here, as well as for the Black Sea, the *GIOP* method application results, presented in the *Level-3m* format, are used. Under negative values of the northern latitude, the corresponding values of southern latitude are meant in Tab. 2.

Results of determining the Pacific Ocean water optical parameters using the *GIOP* method

Date	Coordinates		$a_{ph}(443), m^{-1}$	$a_{dg}(443), m^{-1}$	$C_a, mg \cdot m^{-3}$
	$^{\circ}N$	$^{\circ}W$			
13.06.2010	0.79	135.54	0.046	0.013	0.62
	2.46	135.54	0.015	0.010	0.18
29.05.2003	-1.71	115.29	0.013	0.010	0.16
	-0.71	115.29	0.025	0.012	0.29
	0.87	115.29	0.035	0.013	0.45
	1.12	115.29	0.059	0.010	0.84

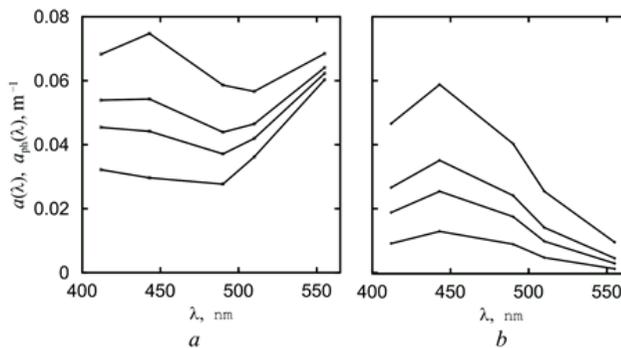


Fig. 4. Results of determining spectral dependencies of the values $a(\lambda)$ – (a) and $a_{ph}(\lambda)$ – (b) based on the Pacific Ocean surveys performed by the SeaWiFS instrument on 29.05.2003

In the considered Pacific Ocean area, there is an interesting phenomenon associated with the El Niño–South Oscillation, accompanied by an occasional increase in the chlorophyll *a* content and a decrease in the water transparency in a narrow band near the equator [21, 22]. Fig. 5 schematically shows the compositional chart of the spatial distribution of the chlorophyll *a* concentration in the equatorial Pacific Ocean relating to the period from May 25 to June 1, 2003, constructed from the results of satellite measurement processing by the traditional method contained in the NASA archive [1].

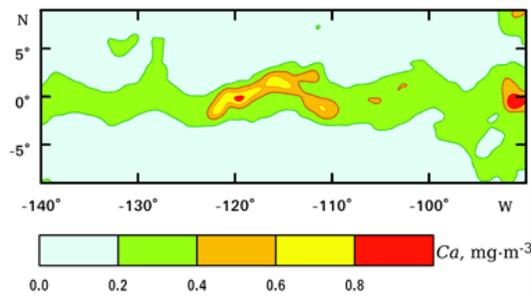


Fig. 5. Compositional chart of C_a distribution in the equatorial Pacific Ocean based on the satellite measurements obtained from May 25, 2003 to June 1, 2003

By the levels of the light absorption coefficient and C_a , the examples in Tab. 2 and in Fig. 4, are well correlated with those discussed above for the Black Sea. Their comparison indicates a significant difference in the laws governing the formation of optical properties of water in the Black Sea and the Pacific Ocean. The main difference is that the values of $a_{dg}(443)$ from Tab. 2 have a relatively low level and almost do not change, and the corresponding variations of $a(443)$ are almost entirely conditioned by $a_{ph}(443)$ variations. At that, the spectra $a(\lambda)$ in the Black Sea and in the ocean are of different shapes (абсолютно нормальное использование).

In a graphic form, the relation between $a_{ph}(443)$ and $a_{dg}(443)$ values in the considered Black Sea and in the Pacific Ocean areas are shown in Fig. 6. Here, in addition to the new data described above, the $a_{ph}(443)$ and $a_{dg}(443)$ determination results, presented in [10, 11], are plotted. In [11], $a_{ph}(443)$ and $a_{dg}(443)$ values from the NASA archive (obtained by *GIOP*) were used, and in [10] they were calculated by a somewhat similar method. At that, for one point, the results of $R_{rs}(\lambda)$ measurements dated 14.09.2000 at the aforementioned station during the expedition in the northeastern Black Sea carried out by spectroradiometer floating on the water surface, were taken as initial data [7]. For the other four points, the $R_{rs}(\lambda)$ spectra, determined according to *SeaWiFS* data, were used in the calculations.

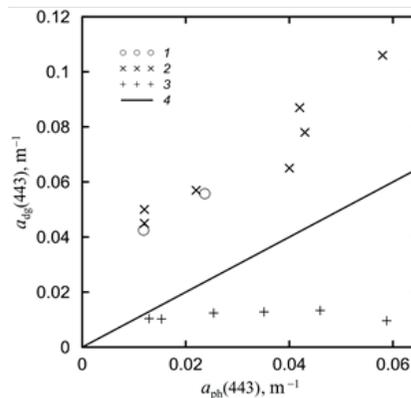


Fig. 6. Correlation between the values $a_{ph}(443)$ and $a_{dg}(443)$: 1 is the Black Sea on 11.08.2006 and on 4.10.2006; 2 is the Black Sea based on the data from [10, p. 82; 11, p. 59]; 3 is the Pacific Ocean on 29.05.2003 and on 13.06.2010; 4 is the straight line $a_{dg}(443) = a_{ph}(443)$

In Fig. 6, there are uniform, but significantly different, trends in each of the two groups of points belonging to different sea areas. It is also interesting that the new data for the Black Sea obtained via *GIOP* are in good agreement with those previously described in [10, 11]. Thus, if the *GIOP* application results for the processing of satellite surveys of the Black Sea and the Pacific Ocean are considered together, the independent variability manifestations of the light absorption components $a_{ph}(\lambda)$ and $a_{dg}(\lambda)$ can be clearly seen. This circumstance plays an important role in the study of the optical properties of seawater, the determination of bioproductivity, etc. If, on the other hand, we confine ourselves only to those data shown in Fig. 6 for the Black Sea, it can be assumed that there is a more or less stable

correlation dependence between $a_{ph}(\lambda)$ and $a_{dg}(\lambda)$, but for its reliable justification a more detailed study, going beyond the scope of this paper, is required.

Conclusions

The analysis performed allows better understanding of the principal possibilities of quantitative interpretation of satellite data in terms of spectral dependences of optically active components of sea water. Although the MODIS and *SeaWiFS* data often contain serious distortions in the Black Sea observations, careful selection of the most reliable data provides useful information on the optical properties of the upper layer of water.

In the examples considered, the spectral dependences of the sea surface reflectances are reproduced with good accuracy by the *GIOP* model used in the operational processing of accumulated satellite measurements. The results of the comparison of the effect of yellow substance dissolved in water and phytoplankton on the total absorption of light in different parts of the spectrum reflect the regional features of the Black Sea and the equatorial Pacific Ocean. In the Black Sea, an important contribution is made by a_{dg} component, whereas a_{dg} plays a weak role in the equatorial Pacific Ocean.

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Vadim S. Suetin – scientific supervision, the problem setting out and formulation, development of mathematical model, the methodology development, qualitative analysis of the results and their interpretation, making conclusions, reference selection and analysis, preparation of the text of the article, writing the summary.

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