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Purpose. The aim of the paper is to study spatial variability of the Mediterranean Sea water optical features as well as to supply the database with additional information on the water optical characteristics required for constructing the regional hydrooptical models. Methods and Results. The data on optical characteristics of the southern Mediterranean Sea waters resulted from the passing measurements in the 2nd cruise of R/V Gorizont in May, 1998 are used in the study. In the water samples from the sea surface layers, the spectral coefficients of beam attenuation and light scattering phase function were measured. Spatial variability of the seawater hydrooptical characteristics from the Strait of Gibraltar to the Dardanelles is considered and analyzed. The equation for the relationship between the asymmetry coefficient of the scattering phase function and the scattering coefficient in the Mediterranean Sea waters is obtained. Conclusions. The data characterizing spatial variability of the southern Mediterranean Sea water optical features in spring are obtained. The coefficients of beam attenuation and scattering in the surface waters decrease gradually from the western Mediterranean Sea to its eastern part. In the Aegean Sea near the Dardanelles, the Marmara Sea waters differing by their high coefficients of beam attenuation and scattering were observed. As for the basic parameters, the scattering phase functions in the Mediterranean Sea waters are similar to those in the Atlantic tropical waters. In the Mediterranean Sea waters as well as in the other water basins, the relationship between the asymmetry coefficient of the light scattering phase function and the scattering coefficient is observed; it is manifested in increase of the phase function asymmetry coefficient with the scattering coefficient.

Keywords: optical characteristics, beam attenuation coefficient, light scattering phase function, phytoplankton, yellow matter.

Acknowledgments: the study was carried out within the framework of the state task on themes No. 0827-2019-0002 and No. 0827-2019-0004.


DOI: 10.22449/1573-160X-2020-1-48-59

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Introduction

Study of the hydrooptical characteristics of the Mediterranean Sea is an actual task in oceanology. This research has been carried out for many years as part of the international programs: POEM (Physical Oceanography of the Eastern Mediterranean), CIESM (The Mediterranean Science Commission, or Commission
Internationale pour l'Exposition Scientifique de la Méditerranée) and MEDCOAST (Mediterranean Coastal Foundation) [1–4].

Since 1972, Marine Hydrophysical Institute (MHI) has been included in the study of the optical properties of Mediterranean waters under the national programs “The World Ocean” and “Satellite Oceanology”. From 1972 to 1992 MHI carried out optical studies in thirteen expeditions in the Mediterranean Sea. A summary of these expeditions is given in [5].

These studies resulted in obtaining extensive data on the optical properties of the Mediterranean Sea waters: spectral beam attenuation coefficient, light scattering phase functions, reflectance band ratio, water column reflectance, Secchi disk depth [5–11]. Most of this data relates to the eastern part of the sea.

In May 1998, on the second cruise of R/V Gorizont from Sevastopol to Lisbon to the international oceanographic exhibition EXPO-98 passing measurements of the optical characteristics of waters in the southern Mediterranean Sea were carried out by Marine Hydrophysical Institute as a part of the Satellite Oceanology program. The research was aimed to study the spatial variability of the optical properties of Mediterranean Sea waters, as well as to supply the database with information on the optical characteristics of water necessary for constructing regional hydro-optical models.

The spectral beam attenuation coefficient and the light scattering phase functions were measurements in water samples taken from the sea surface layers. As a result, data on the optical characteristics of surface waters from the Strait of Gibraltar to the Dardanelles was obtained. The data peculiarity is that it was obtained in a short time and characterize the spatial variability of the optical properties of water in the southern Mediterranean in spring season.

Currently, the relevance of information obtained is determined by the development of regional hydrophysical models, the verification of which is carried out according to field data, including the archival one. This data is also used for reanalysis of hydrophysical fields in assimilation models; it can be used to compare with modern data from optical scanners used for remote sensing of the Mediterranean Sea.

**Instrumentation and Measurement Methodology**

The beam attenuation coefficient (BAC) was measured by a transparency meter [12] in seven spectral regions: 416, 468, 506, 567, 610, 625 and 677 nm (Tab.1). The light scattering phase function was measured with a nephelometer at a wavelength of 520 nm (Tab. 1). Water samples were taken from the sea surface.

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PHYSICAL OCEANOGRAPHY VOL. 27 ISS. 1 (2020) 49
### Technical characteristics of the devices

<table>
<thead>
<tr>
<th>Parameter name</th>
<th>Value</th>
</tr>
</thead>
<tbody>
<tr>
<td><strong>Transparency meter</strong></td>
<td></td>
</tr>
<tr>
<td>Spectral range of BAC measurement, nm</td>
<td>416–677</td>
</tr>
<tr>
<td>Number of spectral channels, pcs.</td>
<td>13</td>
</tr>
<tr>
<td>BAC measurement range of, m(^{-1}) (ln)</td>
<td>0.05–2.0</td>
</tr>
<tr>
<td>BAC measurement Error, %</td>
<td>5</td>
</tr>
</tbody>
</table>

| **Nephelometer**                                    |             |
| Measurement angles of the scattering coefficient \(\sigma(\theta), \theta, ^\circ\) | 2; 7.5 further every 5 up to 162.5 |
| Spectral area of measurements, nm                   | 520         |
| Error of measurements \(\sigma(\theta), \%\)         | 10          |

### Measurement Results and their Discussion

Fig. 1 shows the areas of measurement of optical characteristics along the vessel route. Tab. 2 gives the description of the regions.

*Beam attenuation coefficient.* Fig. 2 shows the BAC value at a wavelength of 416 nm in the measurement regions. A specific feature of spatial variability in the basin from the Alboran Sea to the Levantine Sea is a gradual decrease in the beam attenuation coefficient \(\varepsilon(416)\) from 0.368 to 0.159 m\(^{-1}\). A decrease in \(\varepsilon\) in this part of the basin was also observed at other wavelengths (Tab. 3).

![Fig. 1. Route of the 2nd cruise of R/V Gorizont (crosses with numbers 1–7 denote the regions where hydrooptical characteristics were measured (Table 2))](image-url)

High values of the beam attenuation coefficient \(\varepsilon(416)\) (1.242 m\(^{-1}\)) observed in the northern Aegean Sea near the Dardanelles refers to the waters of the Marmara Sea entering the Aegean through the strait. The waters of the Marmara Sea that emerged from the strait, having lower salinity and, accordingly, lower density, than the waters of the Aegean Sea, spread in its surface layers.
Table 2

The regions where optical characteristics were measured

<table>
<thead>
<tr>
<th>Number</th>
<th>The Mediterranean Sea region</th>
</tr>
</thead>
<tbody>
<tr>
<td>1</td>
<td>The Alboran Sea, the central part</td>
</tr>
<tr>
<td>2</td>
<td>The Balearic Sea, the southern part</td>
</tr>
<tr>
<td>3</td>
<td>The Sicily Strait</td>
</tr>
<tr>
<td>4</td>
<td>The Ionian Sea, the southern part</td>
</tr>
<tr>
<td>5</td>
<td>The Levanteane Sea. Rhodes anticyclone</td>
</tr>
<tr>
<td>6</td>
<td>The Aegean Sea, the middle part</td>
</tr>
<tr>
<td>7</td>
<td>The Aegean Sea near the Dardanelles</td>
</tr>
</tbody>
</table>

Fig. 2. The beam attenuation coefficient in different regions of the Mediterranean Sea, $\varepsilon(416)$, May, 1998 (▲); $\varepsilon$ (the average value in the 0–200 m layer), March–April, 2018 (●)

Table 3

The beam attenuation coefficient ($\varepsilon(\lambda) \cdot 10^3$) in the regions of measurements

<table>
<thead>
<tr>
<th>Wavelength $\lambda$, nm</th>
<th>1</th>
<th>2</th>
<th>3</th>
<th>4</th>
<th>5</th>
<th>6</th>
<th>7</th>
</tr>
</thead>
<tbody>
<tr>
<td>416</td>
<td>368</td>
<td>311</td>
<td>201</td>
<td>253</td>
<td>159</td>
<td>322</td>
<td>1242</td>
</tr>
<tr>
<td>468</td>
<td>299</td>
<td>219</td>
<td>133</td>
<td>196</td>
<td>101</td>
<td>246</td>
<td>1047</td>
</tr>
<tr>
<td>506</td>
<td>276</td>
<td>190</td>
<td>120</td>
<td>179</td>
<td>104</td>
<td>230</td>
<td>1009</td>
</tr>
<tr>
<td>567</td>
<td>311</td>
<td>245</td>
<td>177</td>
<td>223</td>
<td>152</td>
<td>253</td>
<td>1000</td>
</tr>
<tr>
<td>610</td>
<td>478</td>
<td>404</td>
<td>361</td>
<td>391</td>
<td>322</td>
<td>414</td>
<td>1150</td>
</tr>
<tr>
<td>625</td>
<td>529</td>
<td>472</td>
<td>437</td>
<td>467</td>
<td>373</td>
<td>483</td>
<td>1200</td>
</tr>
<tr>
<td>677</td>
<td>644</td>
<td>610</td>
<td>566</td>
<td>598</td>
<td>512</td>
<td>633</td>
<td>1840</td>
</tr>
</tbody>
</table>
In the Mediterranean Sea, a similar spatial variability of the beam attenuation coefficient in the spring period was observed in March – April 2018 according to the data of [3].

The vessel's route [3] passed from the Strait of Gibraltar to the Levant Sea (Fig. 3). Stations with 3000 m depth sounding were carried out along the route. According to averaged data [3, Fig. 2], the vertical distribution of the beam attenuation coefficient in the western and eastern parts of the sea was the same and was characterized by a slight change with depth.

![Fig. 3. The route of the vessel in the Mediterranean Sea in [3] in March – April, 2018 (the dots show location of the stations where the attenuation coefficient was measured)](image)

Tab. 4 shows the mean values of the attenuation coefficient in the 0–200 m layer in some areas according to [3, Tab. 1]. Unfortunately, in the aforementioned work, the wavelength for attenuation coefficient is not indicated. According to the attenuation coefficient values, the measurements were carried out in the red spectrum range, since according to our data the attenuation coefficients at a wavelength of 625 nm (Tab. 3) have similar values. Nevertheless, the spatial variability of the attenuation coefficient given in [3] is of primary importance.

The Tab. 4 data shows a decrease in attenuation in all areas from the Strait of Gibraltar to the eastern Mediterranean (the Ionian Sea and Levantine Sea). As a result, the value of the attenuation coefficient $\varepsilon$ decreased from 0.82 to 0.30 m$^{-1}$.

**Table 4**

**Average value of the attenuation coefficient (\(\varepsilon\)) in the 0–200 m layer in different regions of the Mediterranean Sea based on [3, Table 1]**

<table>
<thead>
<tr>
<th></th>
<th>Average value of the attenuation coefficient, (\varepsilon), m$^{-1}$</th>
</tr>
</thead>
<tbody>
<tr>
<td>The Strait of Gibraltar</td>
<td>0.82 ± 0.10</td>
</tr>
<tr>
<td>The Algerian basin</td>
<td>0.64 ± 0.12</td>
</tr>
<tr>
<td>Corsica</td>
<td>0.53 ± 0.07</td>
</tr>
<tr>
<td>the Sicily</td>
<td>0.48 ± 0.10</td>
</tr>
<tr>
<td>The eastern Mediterranean Sea</td>
<td>0.30 ± 0.09</td>
</tr>
</tbody>
</table>
The Mediterranean Sea is generally characterized by a low content of nutrients and a low level of chlorophyll [13, 14]. The spatial variability of the beam attenuation coefficient in the Mediterranean Sea is explained by the spatial variability of the productivity of surface North Atlantic waters entering it through the Strait of Gibraltar. The content of nutrients in these waters is small [15, 16], and it decreases as the waters move along the southern part of the sea from the Gibraltar Strait to the east due to the consumption of nutrients by phytoplankton. The surface waters are slightly replenished by biogenic matters from deep waters, since the accumulation of nutrients in the deep waters of the sea is hindered by the continuous outflow of these waters back to the Atlantic Ocean through the Strait of Gibraltar [15]. Vertical transport exists in limited frontal and few upwelling areas [14]. Due to these factors, the biological productivity of surface sea water and, accordingly, the concentration of phytoplankton in them decrease, which leads to an increase of the water transparency [17, p. 229–234].

The above facts does not apply to the Aegean Sea waters, since the North Atlantic waters do not enter this sea.

Tab. 3 and Fig. 4 show the attenuation coefficient spectra ε(λ) in different regions of the Mediterranean Sea. The spectrum shape in the Aegean Sea waters near the Dardanelles, where the waters of the Marmara Sea were registered, differs by the position of the minimum ε(λ)_{min} – in the Mediterranean Sea it is at wavelengths of 468–506 nm, in the waters of the Marmara Sea – at 567 nm. Such a shift in the minimum in the attenuation coefficient spectrum occurs at high concentrations of dissolved organic matter (yellow matter) in water. The absorption coefficient of yellow matter is maximal in the short-wave spectrum region and decreases with the law as κ(λ)_{ym}~e^{μλ} with increasing wavelength. With an increase in the yellow matter concentration, the effect of κ(λ)_{ym} on the increase of attenuation coefficient reaches longer wavelengths, which leads to a shift of the minimum ε(λ) to the long wavelength region.
According to [18], the amount of yellow matter in the Mediterranean Sea is twice its content in the nearest Atlantic sector. In addition, the values of absorption coefficient $\kappa(442)_{\text{ym}}$ by dissolved organic matter is almost 50% higher in the western part of the sea than in the eastern part. Spectra 1–5 (Fig. 4) reflect the indicated trend: due to an increase in the absorption coefficient of yellow matter in the direction from east to west of the Mediterranean Sea, the beam attenuation coefficient at 416 nm wavelength increases accordingly.

**Light scattering phase functions.** Fig. 5 shows the beam scattering coefficients, asymmetry coefficients, and average cosines of scattering in different regions. The spatial variability of the scattering coefficients is similar to the attenuation coefficient variability (Fig. 2). The reasons for this variability are discussed above.

### Table 5

Parameters of the light scattering phase functions

<table>
<thead>
<tr>
<th>Region number</th>
<th>Scattering coefficient $\sigma$, 1/m</th>
<th>Asymmetry coefficient $K$</th>
<th>Average cosine of scattering, $\cos \theta$</th>
</tr>
</thead>
<tbody>
<tr>
<td>1</td>
<td>0.19</td>
<td>58.7</td>
<td>0.946</td>
</tr>
<tr>
<td>2</td>
<td>0.21</td>
<td>77.8</td>
<td>0.960</td>
</tr>
<tr>
<td>3</td>
<td>0.09</td>
<td>33.9</td>
<td>0.919</td>
</tr>
<tr>
<td>4</td>
<td>0.15</td>
<td>56.0</td>
<td>0.944</td>
</tr>
<tr>
<td>5</td>
<td>0.11</td>
<td>33.9</td>
<td>0.916</td>
</tr>
<tr>
<td>6</td>
<td>0.17</td>
<td>54.6</td>
<td>0.944</td>
</tr>
<tr>
<td>7</td>
<td>0.27</td>
<td>38.5</td>
<td>0.906</td>
</tr>
</tbody>
</table>

Tab. 6 shows the minimum and maximum values of the phase function parameters observed in the Mediterranean Sea, and for comparison, similar phase function parameters in the Tropical Atlantic waters [19]. The comparison shows the proximity of the values of all parameters.

### Table 6

Range of the values of the phase function parameters observed in the Mediterranean Sea and the Tropical Atlantic [15]

<table>
<thead>
<tr>
<th>Region</th>
<th>Scattering coefficient $\sigma$, 1/m</th>
<th>Asymmetry coefficient $K$</th>
<th>Average cosine of scattering, $\cos \theta$,</th>
</tr>
</thead>
<tbody>
<tr>
<td>the Mediterranean Sea</td>
<td>0.09–0.21</td>
<td>34–78</td>
<td>0.916–0.960</td>
</tr>
<tr>
<td>the Tropical Atlantic</td>
<td>0.09–0.25</td>
<td>34–77</td>
<td>0.927–0.962</td>
</tr>
</tbody>
</table>
In the Mediterranean waters, a close relationship was observed between the asymmetry coefficient of the phase function and the scattering coefficient $K = f(\sigma)$ (Fig. 5). Such relations in water basins are regional and depend on the composition of the suspended matters in water [20]. In Fig. 5 shows the point for the Aegean waters near the Dardanelles, falling out of the $K = f(\sigma)$ relationship for the Mediterranean Sea. These parameter values refer to the waters of the Marmara Sea entering the Aegean Sea through the strait. The position of this point on the graph (Fig. 5) according to the model developed in [20] indicates a high concentration of fine suspended matter in the waters of the Marmara Sea.

![Graph showing relationship between asymmetry coefficient of scattering phase function and scattering coefficient](image1)

**Fig. 5.** Relationship between the asymmetry coefficient of scattering phase function $K$ and the scattering coefficient $\sigma$: ◆ – the Mediterranean Sea, ▲ – the Marmara Sea

Fig. 6 and Tab. 7 show examples of phase functions in the Mediterranean Sea and in the waters of the Marmara Sea.

![Graph showing light scattering phase functions](image2)

**Fig. 6.** Light scattering phase functions in the Mediterranean Sea: ▲ – the Alboran Sea (region 1); ■ – the Levantine Sea (region 5), ● – the Marmara Sea (region 7)
The light scattering phase functions in the Mediterranean Sea waters

<table>
<thead>
<tr>
<th>Scattering angle $\theta$, °</th>
<th>Ln $\sigma(\theta)$, m$^{-1}$</th>
<th>in region 1</th>
<th>in region 5</th>
<th>in region 7</th>
</tr>
</thead>
<tbody>
<tr>
<td>2.0</td>
<td>1.334</td>
<td>0.713</td>
<td>1.54</td>
<td></td>
</tr>
<tr>
<td>7.5</td>
<td>-1.863</td>
<td>-2.852</td>
<td>-2.23</td>
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<tr>
<td>12.5</td>
<td>-3.404</td>
<td>-3.795</td>
<td>-2.69</td>
<td></td>
</tr>
<tr>
<td>17.5</td>
<td>-4.094</td>
<td>-4.347</td>
<td>-3.22</td>
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<tr>
<td>22.5</td>
<td>-4.692</td>
<td>-5.083</td>
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<tr>
<td>27.5</td>
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<td>-5.505</td>
<td>-4.35</td>
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<tr>
<td>32.5</td>
<td>-5.819</td>
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</tr>
<tr>
<td>37.5</td>
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<td>-4.99</td>
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</tr>
<tr>
<td>42.5</td>
<td>-6.486</td>
<td>-6.877</td>
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</tr>
<tr>
<td>47.5</td>
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<td>-7.061</td>
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<td>62.5</td>
<td>-7.337</td>
<td>-7.613</td>
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</tr>
<tr>
<td>67.5</td>
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<td>-7.682</td>
<td>-6.56</td>
<td></td>
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<td>-7.659</td>
<td>-7.958</td>
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<td>77.5</td>
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<td>-8.234</td>
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<td>122.5</td>
<td>-8.510</td>
<td>-8.510</td>
<td>-7.84</td>
<td></td>
</tr>
<tr>
<td>127.5</td>
<td>-8.648</td>
<td>-8.510</td>
<td>-7.70</td>
<td></td>
</tr>
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<td>-8.487</td>
<td>-8.625</td>
<td>-7.68</td>
<td></td>
</tr>
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<td>137.5</td>
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<td>-8.648</td>
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<tr>
<td>142.5</td>
<td>-8.556</td>
<td>-8.556</td>
<td>-7.82</td>
<td></td>
</tr>
<tr>
<td>147.5</td>
<td>-8.625</td>
<td>-8.464</td>
<td>-7.73</td>
<td></td>
</tr>
</tbody>
</table>
Conclusion

In May 1998, data characterizing spatial variability of the optical properties of the waters of the southern Mediterranean Sea in spring season were obtained from the measurements taken on R/V Gorizont.

The coefficients of beam attenuation and scattering in surface waters are gradually decreasing in the direction from the western Mediterranean Sea to the eastern one. The same spatial variability of the attenuation coefficient in the Mediterranean Sea in the spring (March – April 2018) is observed according to field data in [3].

In the Aegean Sea near the Dardanelles, the Marmara Sea waters are observed, characterized by high coefficients of the beam attenuation and scattering.

The parameters of light scattering phase functions in the Mediterranean Sea waters are similar to those in the Atlantic tropical waters.

In the Mediterranean Sea waters, as in other water basins, there is a relationship between the asymmetry coefficient of the light scattering phase function and scattering coefficient, which is manifested in increase of the phase function asymmetry coefficient with the scattering coefficient.

Presented data on the optical characteristics of the surface Mediterranean Sea waters can later be used for a comparative analysis of empirical data on the optics of Mediterranean waters obtained in the framework of field research and remote sensing of the Mediterranean Sea.

REFERENCES


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*Contribution of the co-authors:

Viktor I. Mankovsky – the problem statement, preparation of the article text, processing, interpretation and description of the study results, formulation of conclusions

Ekaterina V. Mankovskaya – selection and analysis of literature, presentation of data in the text and their analysis, preparation of graphic and text materials, article correction

All the authors have read and approved the final manuscript.

The authors declare that they have no conflict of interest.